

Morphometric analysis of the Catchments of Himayatsagar and Osmansagar Reservoirs - Hyderabad

V.Varalakshmi

Women Scientist, DST, Department of civil engineering
Marri Laxman Reddy Institute of technology and management, Dundigal, Hyderabad, India
varasays@yahoo.co.in

Abstract— The linking of the geomorphological parameters with the hydrological characteristics of the basin provides a simple way to understand the hydrologic behaviour of the different basins particularly of the ungauged basins because the hydrologic and geomorphic effects of natural and human process with in a catchment are focused at its outlet. The development of morphometric techniques was a major advance in the quantitative description of the geometry of the drainage basins and its network which helps in characterizing the drainage network. The geomorphological properties which are important from the hydrological studies point of view include the linear, aerial and relief aspect of the watersheds. In the present study morphological study has been carried in the catchments of Himayatsagar and Osmansagar reservoirs. The catchments of Himayatsagar and osmansagar reservoirs cover the area of 1315.14 and 748 km within the elevation rang of 575 to 715 and 585 to 705 m resepectively.the drainage pateern varies from dentric to subdentric. They are 5th and 4th order drainage basins respectively. The stream length and stream areas to stream number relation follows the Hortons laws. Stream density shows low value for all stream orders. The bifurcation ratio is high for Osmansagar reservoir catchment. This is due to the variation in stream frequencies in successive orders. Lower value of drainage density and stream frequency indicates the region of highly permeable subsoil strata and the catchment will not generate runoff soon. The length of overland flow and time of concentration for the catchments of Himaytsagar and osmansagar is estimated to be .68, .70 and 16.74, 15.81hours resepectively. From the values of Circularity ratio, elongation ratio, form factor, and compactness coefficient of the catchments of Himayatsagar and Osmansagar reservoirs it is expected that the catchment is highly irregular, strong relief, steep ground slope and elongated hydrograph. The high ruggedness number of the catchment of Himayatsagar and Osmansagar reservoir shows the high roughness or unevenness.

Keywords— *Drainage network, Morphometric Analysis*

1. Introduction

Hyderabad is the fifth largest city in India is facing with

severe water shortage on account of growing population with the rate of more than 5% and depleting water resources. At present, only one-half hour to two hours every alternate day in MCH area and one hour every alternate day in surrounding municipalities of Hyderabad and in summer even this little supply is in alternate days. For over a 6.8 million population of the city, the actual requirement is at the rate of 290 Million gallons per Day (Mgd) but the present water availability is 270Mgd (Hyderabad city plan).

Himayatsagar and Osmansagar are the two oldest reservoirs supplying the water to the city apart from Manjira and singur dams. Over the years the inflows to the Osmansagar and Himayatsagar are slowly decreasing in spite of the average annual rainfall is more or less the same for the past 46years (Dr.B.Venkateswara Rao, 2009). This shows the importance of Hydrogeological studies in these areas.

2. Location And Description of the Study Area

The study area for the present work consists of catchment areas of Himayatsagar and Osmansagar reservoirs (Fig.1). The Himayatsagar reservoir was constructed on Esa river in 1925 and is located 9.6 km in southwest direction from the city and Osmansagar reservoir was constructed on Musi river in 1922, and is located 9.6 k.m. from Hyderabad in western direction. The Himayatsagar catchment covers 153 villages and spread over 1315 km² area with a slope range of 575 m to 715 m Where the Osmansagar catchment area consists of 748 km² with a slope varies between 585 m to 705 m and covers 93 villages. The two reservoirs are supplying drinking water to Hyderabad city, which is extending from 17^o 10' to 17^o50' of North latitude and 78^o 10' to 78^o50' of East longitude.

Both the catchments of Himayatsagar and Osmansagar reservoirsis shows flat to gently undulating topography except for few hillocks and valleys. The region is mainly underlain by a peninsular gneissic complex that includes a variety of granites, magmatites of various phases and enclaves of older metamorphic rocks belonging to the Archean age. These are intruded by various acidic (pegmatite's, aptite, quartz veins/reefs) and basic intrusives of dolerite and gabbros. Major dykes of dolerite composition cut across the country rocks in different

directions. The predominant soils in the basin are sandy loam, clay loam, black cotton soils and rocky soils. The riverbed is mostly deposited with sandy soils. Ground water occurs under Unconfined to Semi-confined conditions in weathered and fractured formations respectively. Three seasons prevail in the basin: summer (March – May), monsoon/ rainy (June – November) and Winter (December – February). The mean annual rainfall is around 800 mm. 75% of rainfall occurs during the monsoon period. The minimum and maximum temperatures vary between 12⁰ C and 43⁰C.

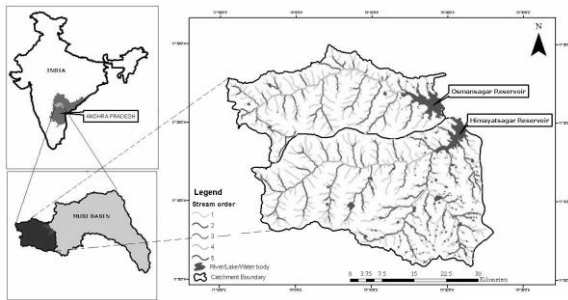


Fig.1 Location map of the study area

3. Literature Review

The development of morphometric techniques was a major advance in the quantitative description of the geometry of the drainage basins and its network. These parameters are useful in characterizing river basins and comparing their characteristics. For the first time it was proposed by Horton (1945). Later it was modified by Strahler (1957). Naithani & Rawat the Morphometry parameters could be categorized into two major groups namely; the measured parameters such as basin area, axial length etc. and the calculated parameters like drainage density, stream frequency, etc. Sandeera et al (2007). studied the quantitative description of the geometry of the Kotmale reservoir and its network and compared the characteristic of several drainage networks and examining the effect of variables such as lithology, rock structure, rainfall etc. Yasuo Shimano (2005) analyzes the composition of stream networks of 180 drainage basins in the entire country of Japan and discuss the characteristics of drainage basins in the Japanese Islands. P.D.Sreedevi (2004) carried out morphometric analysis for obtaining groundwater potential zones in a structurally controlled terrain of total area of the Pageru River basin.

4. Methodology

To estimate the morphometric features of catchments of Himayatsagar and osmansagar reservoirs, the drainage network is drawn with the digital elevation model in arc GIS software by using the toposheet of 1:2, 50,000 scale.

Stream order: The first step in drainage basin analysis is to designate the stream order. Stream order is introduced by Horton. Later it is modified by Strahler. The smallest streams of the network, which have no tributaries, are called first order streams. When two first order streams coalesce they form a second order stream and further along its course this stream may join another second order channel to form one of the third order and so on. A lower order stream, such as one of the first order joining another higher order does not alter the rank of the latter. The relevant numbers were entered into the attribute table of the drainage network using GIS software.

Axial length of the basin (Lb km): It is the distance from the outlet to the most remote point on the basin

Mean length (Lu km): Mean length of channel Lu of order U the total length is divided by the number of segments Nu of that order

$$Lu = \sum Lu / Nu$$

Stream length ratio (Rl): It is the ratio of the meanlength of segments of order U to the mean length segment of the next lower order.

$$Rl = Lu / Lu - 1$$

Drainage basin area: Drainage basin area Au is defined as the total area projected on a horizontal plane contributing overland flow to the channel segments of the given order including all the lower orders. The stream areas or derived by using DEM model in GIS software.

$$Au = (\sum A1 + \sum A2 + \sum A3 \dots \dots \dots + \sum Ao1) + \sum Ao2 + \sum Ao3 \dots \dots \dots + \sum Aou$$

Drainage density: Drainage density Dd is introduced by Horton. It is the ratio of the total length of all stream channels in the basin to the total area of the basin.

$$Dd = \sum Lu / Au$$

Constant channel maintainace (Cc): According to Schumm Constant channel maintainace is the inverse of the drainage density.

$$Cc = 1 / Dd$$

Stream frequency (Fc): Horton introduced stream frequency Ds as the number of stream segments per unit area.

$$Ds = \sum Nu / Ak$$

Length of overland flow (Lg): According to Horton definition length of overland flow approximately equal to half of the reciprocal of the drainage density.

$$Lg = 1 / 2 Dd$$

Time of Overland flow: This equation is suggested by Kirpich for time of overland flow is

$$T_c = 0.0195L^{0.77}S^{-0.385}$$

Where T_c is in minutes, L is the maximum length of travel of water along the water course in meters and S is the slope expressed as the ratio of the difference in elevation between the remotest point and the catchment outlet to the length L .

Circularity ratio (R_c): According to Miller circularity ratio R_c defined as the ratio of basin area A_u to the area of circle A_c having the same perimeter as the basin.

$$R_c = A_u/A_c$$

Elongation ratio (R_e): Schumm defined the elongation ratio R_e as the ratio of diameter of a circle of the same area as the basin D_c to the maximum basin length L_b

$$R_e = D_c/L_b$$

Form factor (F_f): Horton defined the form factor F_f which is the ratio of basin area A_u to the square of basin length L_b .

$$F_f = A_u/L_b^2$$

Compactness factor (C_c): The compactness factor was obtained from the ratio of the perimeter of the watershed P to the total drainage basin area. (Gupta, 1999)

$$C_c = P/2\sqrt{A}$$

Laminate ratio: It is the ratio of square of the length of the basin to the four times of the basin area

$$L_b^2/4A_u$$

Maximum basin relief (R): It is defined as the height difference between the maximum elevation on the watershed Z_m and the elevation of the Gauging station Z_o

$$R = Z_m - Z_o$$

Relief ratio (R_h): Schumm measured relief ratio R_h as the maximum basin relief to horizontal distance along the longest dimension of the basin parallel to the principle drainage line.

$$R_h = R/L_b$$

Relative relief (R_{hp}): It is the ratio of the maximum basin relief to the perimeter of the basin (Shivapur, A.V., 2006)

$$R_{hp} = R/P$$

Ruggedness number (R_n): This is the product of Drainage density and Catchment relief.

$$R_n = D_d * R$$

5. Results and Discussions

The main or trunk stream and its tributary streams that drain the basin area collectively form the stream network. The spatial arrangement of a river and its tributary streams in a drainage network of Himayatsagar reservoir catchment and Osmansagar reservoir catchment is referred to as dendritic too sub-dendritic drainage pattern of a basin. Figure 1 illustrates the stream network of catchments of Himayatsagar reservoir and Osmansara reservoir, the stream orders of Himayatsagar reservoir catchment order, varied from 1 to 5 and the total number of stream segments of all orders recorded was 423. stream network indicated 335 of 1st orders, 72 of 2nd Orders of 12 of 3rd order, 3 of 4th order and 1 of 5th order streams where The catchment of osmansagar reservoir stream orders varied from 1 to 4 and the total number of stream segments of all orders recorded was 274. Orders of stream network indicated 219 of 1st order, 44 of 2nd order, 10 of 3rd order and 1 of 4th order streams.

The mean length and mean areas of the streams shows an increasing trend with stream order (Table 1 and table.2). The number of streams, stream order and stream areas follows the Hortons laws, i.e, it has been observed that the logarithm of number of streams decreases with increasing stream order. This is Hortons I law. Logarithm of the mean length of stream increases with increasing stream order this is Horton II law. Logarithm of mean area of stream increase with increasing stream order. This is Horton III law.

According to Gregory & Walling (1985) almost every watershed characteristics is correlated with the catchment area. Kale and Gupta (2001) stated that the larger basins have large average discharge. The areas of the catchments

Table.1 Stream parameters of the Catchment of Himaytsagar Reservoir

Stream order(U)	No. of streams (Nu)	Length of Streams (Lu)	Areas of the Streams (Au)	Log Mean stream length (logLu)	Log Mean stream area (LogAu)	bifurcation ratio (R _b)	strem length ratio (R _L)	Stream density
1	335	554.4406	911.79	0.21	.43	4.652778	1.9456198	.60
2	72	231.8466	1048.92	0.50	1.16			.74
3	12	103.41	1226.92	0.93	2.00			.72
4	3	50.71596	1284.28	1.22	2.63			.732
5	1	21.87	1315.15	1.34	3.11			.731

Table.2 Stream parameters of the Catchment of the Osmansagar Reservoir

Stream order(U)	No. of streams (Nu)	Length of Streams (Lu)	Areas of the Streams (Au)	Log Mean stream length (logLu)	Log Mean stream area (LogAu)	bifurcation ratio (R _b)	stream length ratio (R _L)	Stream density
1	219	280.86	483.34	.10	.34	4.97 4.40 10.00	.49 .40 .99	.58
2	44	140.03	602.8	.50	1.13			.69
3	10	56.02	633.49	.74	1.80			.75
4	1	55.90	748.14	1.74	2.87			.71

of Himayatsagar and Osmansagar are 1316 km² and 748.14 km² respectively.

The length and slope of the catchment effects the lag time taken for the water to reach the reservoir from its longest distance in the catchment after the rainfall. The length and maximum basin relief of the catchments of Himayatsagar and Osmansagar are 58 km, 140 m and 53.8 km, 120 m respectively. From Fig.2 it shows that there is a decrease in slope with the length along the stream towards mouth of the drainage basin.

The bifurcation ratios of the Himayatsagar reservoir catchment varies between 3 to 6. This shows that the geologic structures with in the catchment do not distort the drainage pattern. It also indicates that a basin is a mature topography due to the result of the process of drainage integration. Where the osmansagar reservoir catchment varies between 4 to 10. These high bifurcation ratios are the result of large variation in stream frequencies in successive orders and The structural control encourages the development of elongate narrow drainage basins.

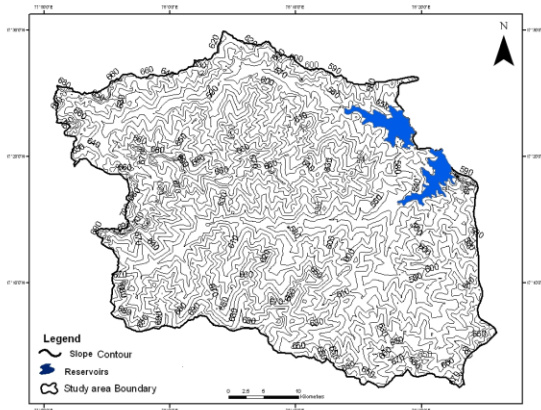


Fig.2 Slope Map of the Study Area

The drainage density is indicative of the closeness of spacing of channels thus providing a quantitative measure of average length of stream channel for the whole basin. It also indicates about the physical properties of the underlying rocks. The drainage densities of the catchments of the Himayatsagar and osmansagar is .731 km per km²

and .71 km per km² indicates region of highly subsoil strata and coarse texture (Smith,1950 and Strahler, 1957). This coarse texture gives more time for overland flow. The stream frequency also shows the low value of .321 number of streams per km² and 0.36 number of segments per km². This shows the catchment will not generate runoff soon and low sediment production. This is evident from the values of length and time of overland flow which is 1.36 km and 16.7 hours.

We know that the pattern of flow hydrograph and its time variance depends on shape of the basin flow response to catchment precipitation and its topography. The important parameters that describe the shape of the basin are circularity ratio (R_c), elongation ration (R_e) compactness coefficient and form factor (F_f). From the values of Circularity ratio, elongation ratio, form factor, and compactness coefficient of the catchments of Himayatsagar and Osmansagar reservoirs it is expected that the catchment is highly irregular, strong relief, steep ground slope and elongated hydrograph (Table.3). It also shows the basin is mature and old. It is evident from the relief ratio value of the catchment. From The high ruggedness number of the catchment of Himayatsagar and Osmansagar reservoir shows the high roughness or unevenness. High relative relief shows high region or low relative relief shows paddy plain and valley.

6. Conclusions

The catchments of Himayatsagar and osmansagar reservoirs cover the area of 1315.14 and 748 km within the elevation rang of 575 to 715 and 585 to 705 m respectively. the drainage pateern varies from dentic to subdentic. They are 5th and 4th order drainage basins respectively. The stream length and stream areas to stream number relation follows the Hortons laws. Stream density shows low value for all stream orders. The bifurcation ratio is high for Osmansagar reservoir catchment. This is due to the variation in stream frequencies in successive orders. Lower value of drainage density and stream frequency indicates the region of highly permeable subsoil strata and

Table.3 Morphometric Parameters of the Catchments of Himayatsagar and Osmansagar Reservoirs

S.No.	Paraameter	Himayatsagar	Osmansagar
1	Area of catchment in km ²	1315.70	748.14
2	Perimeter in km	174.69	151.06
3	Axial length in km	58.02	53.80
4	Breadth in km	30.83	19.97
5	Drainage density	.73	0.71
6	Constant channel maintainance	1.36	1.40
7	Stream frequency	.32	0.36
8	Length of overland flow	.68	0.70
9	Time of overland flow (hr)	16.74	15.81
10	Circularity ratio	.54	0.41
11	Elongaton ratio	.70	0.57
12	Form factor	.39	0.25
13	Compactness factor	1.35	1.55
14	Lminscate ratio	.63	.96
15	Max water shed relief	140	120
16	Relief ratio	2.41	2.23
17	Relative relief	.80	0.79
18	Ruggedness number	102.39	85.46

the catchment will not generate runoff soon. The length of overland flow and time of concentration for the catchments of Himaytsagar and osmansagar is estimated to be .68, .70 and 16.74, 15.81hours respectively. From the values of Circularity ratio, elongation ratio, form factor, and compactness coefficient of the catchments of Himayatsagar and Osmansagar reservoirs it is expected that the catchment is highly irregular, strong relief, steep ground slope and elongated hydrograph. The high ruggedness number of the catchment of Himayatsagar and Osmansagar reservoir shows the high roughness or unevenness.

References

- [1] Gregory KJ, Walling DE., Drainage basin form and process: a geomorphological approach. Wiley, New York, 456 pp,1973
- [2] Horton R., Erosional development of streams and their drainage basins. Hydrophysical approach to quantitative morphology. Geol Soc Am Bull 56(3):275–370,1945
- [3] Kale, V. S. and Gupta, A., Introduction to Geomorphology, Orient Longman, Calcutta.,2001
- [4] Sreedevi PD, Subrahmanyam K, Ahmed S.,The significance of morphometric analysis for obtaining groundwater potential zones in a structurally controlled terrain. Environ Geol 47:412–420,2004
- [5] Strahler AN., Quantitative analysis of watershed geomorphology. Trans Am Geophys Union 38:913–920.,1957
- [6] Yasuo Shimano., Hydraulic Geometry of the Drainage Basin of Iruma River”, Jour. Of Hydrology Vol.4, PP.20-24, 1971
- [7] Venkateswara Rao.B, Varalakshmi.V and Vijayasarada S.T., Ground Water Recharge in Hard Rock Areas of Musi Basin, Journal of Applied Hydrology, Vol.No.XXII, No.1, pp102-117,2009



Dr.V.Varalakshmi is a principle investigator of DST project , Department. of civil Engineering, Marri Laxman Reddy Institute of Technology and Management. She holds his Ph.D from Jawaharlal Nehru Technological University Hyderabad and Master Degree from Andhra University. She published more than 30 papers in various national

and international journals and conferences and having nearly 10 years teaching and research experience